

## Transport of Pollutants in the Atmosphere

The atmosphere is the environmental medium where we live and breath. Modeling of atmospheric pollution can be used to determine human exposure to existing pollution sources and to predict future exposures from industrial accidents. There are many sources of atmospheric pollution, including volcanoes, industrial smoke stacks, fugitive (or nonpoint) industrial emissions, gasoline stations, forest fires, industrial accidents, and automotive and railroad accidents. In Fate<sup>®</sup>, we will develop relatively simple models to predict the fate and transport of pollution released such sources.

First, we will compare other fate and transport models to the general atmospheric model. The aquatic models in Fate<sup>®</sup> were given only for one or two dimensions. Streams and lakes can be adequately modeled using one-dimensional models since most of the dispersion is in the longitudinal direction, while groundwater systems require at least two dimensions (x and y). Two dimensions are required in the latter system because the groundwater is not constrained by a river or lake bank, and dispersion can occur in all directions. Vertical dispersion, while important near a point pollution source, becomes less important when the groundwater system is bounded by confining layers above and below the aquifer of interest, which is why we used the simpler two-dimensional model in the instantaneous and pulse groundwater releases.

While the aquatic models may have seemed complicated, they are simpler than most atmospheric models. Because of wind currents and mixing, atmospheric models have to incorporate three dimensions, which automatically makes the governing equations more complex. As usual, we make many assumptions that make our model more manageable. For example, the models given in Fate<sup>®</sup> are not designed for gases that are more or less dense than the atmosphere, and therefore ignore buoyancy effects. The models distinguish between step and instantaneous sources, although actual atmospheric pollution episodes can lie between these two extremes. Unlike the aquatic models that allow first-order decay processes, our atmospheric models do not allow degradation of pollutants. This assumption is justified for models of a pollutant over relatively short distances (under 10 000 meters or 7 miles) because most photochemical reactions (except for the production of smog) require the pollutant to be in the atmosphere over a much longer time frame (hours to days). The dominant force resulting in the reduction of the pollutant concentration is dispersion, which can rapidly dilute pollutant concentrations. However, understanding and accounting for dispersion can be very complicated. First, we will look at the movement of atmospheric gases over the Earth's surface.

A profile of the wind's velocity with increasing height has a steep increasing parabolic shape, with low velocity at the Earth's surface due to friction between the moving air and the ground. The surface wind velocity, is also subject to many complex variable, however. For example, the roughness of the Earth's surface can significantly impact the shape or steepness of the wind velocity-height profile. The wind velocity profile over an open grassland is illustrated in right-hand side of Figure 1, showing that wind speed rapidly approaches its maximum as height above the surface increases.

Compare this to an urban setting where tall buildings impede the path of the wind and slow its speed. This expands the velocity-height gradient well above the Earth's surface. The resulting lower wind velocity could decrease the turbulence and subsequent dispersion by slowing the wind velocity but may also result in stagnant pockets of the atmosphere that can contain clear or polluted air. Thus, the increase in the surface's roughness from the presence of building will greatly affect flow patterns and ground-level pollutant concentrations. Variables such as this demonstrate that, atmospheric processes are too complicated even for our most sophisticated models. In our brief introduction to we will simplify our model by assuming that an average wind speed can be used and, in general, we will not account for differences in surface roughness.

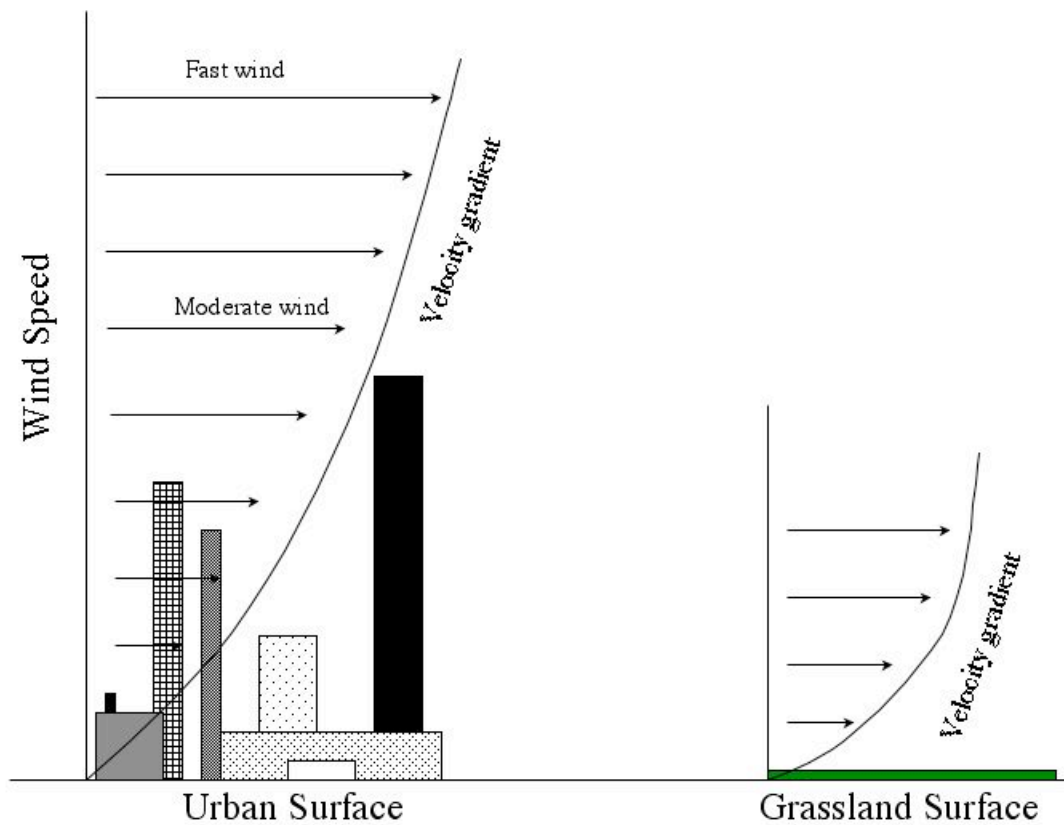


Figure 1. The effect of surface roughness on wind speed.

While surface roughness can greatly affect turbulence and mixing, the magnitude of wind speed can also increase mixing. We will refer to this mixing as dispersion, since the net result is a dilution of pollutant concentrations. If we combine the effects of wind velocity and the atmospheric temperature as a function of height above the surface, we obtain the three basic turbulence scenarios shown in Figure 2. We will start with an isolated pocket of atmosphere at nighttime temperatures (shown in Figure 2a). This type of condition occurs where a thick cloud layer prevents the Earth from radiating its heat to space as it cools during the night. Under these conditions, an emission from an industrial stack will take the shape of the plume shown in Figure 2a. The released gases

will rise or sink until their density (temperature) matches that of the surrounding (diluting) atmospheric gases. Then the plume will take the shape of a thin layer.

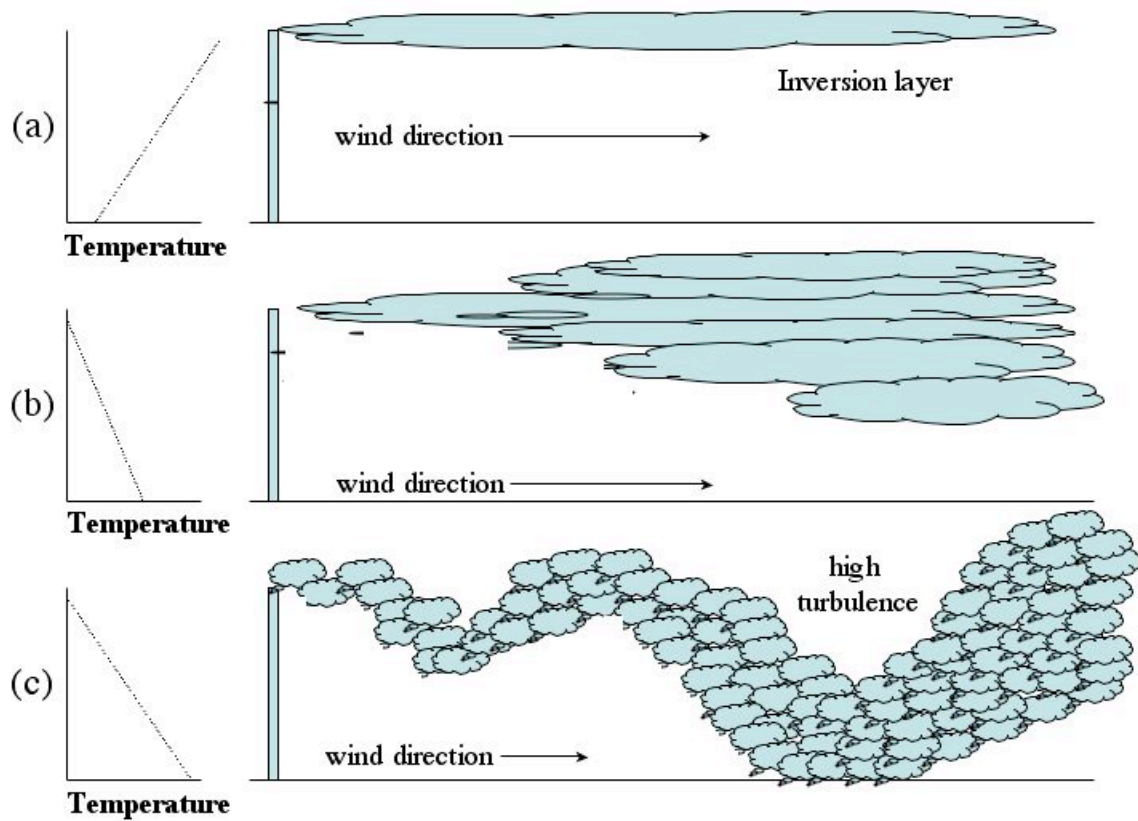


Figure 2. Three basic turbulence scenarios for plumes.

Under daytime heating conditions, the temperature-height profile will be similar to the one shown in Figure 2b. In a steady wind, the plume will spread in all directions, but mostly in the longitudinal direction. With a lower temperature-height gradient and a higher wind velocity, extreme turbulence will be observed (Figure 2c). In order to attempt the modeling of these conditions, we must greatly simplify the temperature and wind relationships.

We will start our simplification process by attempting to combine the effects of wind velocity, temperature-height profiles, and cloud cover into a set of atmospheric stability categories. As we do this, remember that our goal is to come up with a way to characterize dispersion (mixing) of the pollutant with the atmospheric gases. Table 1 shows a qualitative approach to the combined effects of wind speed and cloud cover collected for rural settings in England. Cloud cover is a good reflection of heat back to the Earth. The categories range from strongly unstable (category A reflected in Figure 2c) to very stable (category G) and distinguish between day and night conditions.

Table 1. Pasquill Stability Categories

| Wind speed<br>(at 10 m<br>elevation<br>m/s) | Day, Degree of Cloud Insolation |          |        | Night  |                                 |
|---|---------------------------------|----------|--------|--|---------------------------------|
|   | Strong                          | Moderate | Slight | Thinly<br>overcast or<br>greater than<br>50% low<br>clouds | Less than<br>50% cloud<br>cover |
| < 2   | A                               | A-B      | B      | G  | G                               |
| 2 – 3                                       | A-B                             | B        | C      | E  | F                               |
| 3 – 5                                       | B                               | B-C      | D      | D  | E                               |
| 5 – 6                                       | C                               | C-D      | D      | D  | D                               |
| >6  | C                               | D        | D      | D  | D                               |
|   |                                 |          |        |  |                                 |

Source: Turner (1994) and Pasquill (1961)

Turner (1994) adds the following notes on selecting the appropriate category:

1. Strong insolation corresponds to sunny midday in midsummer in England; slight insolation to similar conditions in midwinter.
2. Night refers to the period from 1 hour before sunset to 1 hour after sunrise.
3. The neutral category D should also be used, regardless of wind speed, for overcast conditions during day or night and for any sky condition during the hour preceding or following night as defined in note number 2.

Next, the somewhat qualitative categories in Table 1 are used to mathematically predict values for horizontal dispersion coefficients (Table 2), which are estimates of mixing in the x and y directions. We do not have a way to accurately mathematically predict these values, and the data in Tables 1 and 2 are empirical (based on experimental observations). We usually assume that dispersion in the x and y directions is the same; thus Table 2 can be used to estimate  $\sigma_x$  and  $\sigma_y$  simultaneously. The equations given in Table 1 were used to draw the lines in Figure 3. Note that dispersion increases as you move away from the point source of pollution. This should be intuitive since mixing continues and the wind causes more mixing as you move away from the point source and the wind will also cause more mixing as you move away from the source. So, for every pollutant concentration you attempt to estimate, you must select a distance from the point source. The unfortunate result of this is that Fate<sup>®</sup> can only plot a slice of the concentration in the y and z planes. You will have to manually plot the concentration gradient in the x, or longitudinal, direction.

Table 2. Pasquill-Gifford Horizontal Dispersion Parameters (Turner, 1994)

| $\sigma_y = 1000 * \tan(T)/2.15$   |                              |
|--|------------------------------|
| <p>where x is the downwind distance (in km) from the point source and T is one-half Pasquill's <math>\sigma</math> in degrees. T, as a function of x, is determined by each stability category from Table 1.</p> |                              |
| Stability  | Equation for T               |
| A  | $T = 24.167 - 2.5334 \ln(x)$ |

|   |                              |
|---|------------------------------|
| B | $T = 18.333 - 1.8096 \ln(x)$ |
| C | $T = 12.5 - 1.0857 \ln(x)$   |
| D | $T = 8.333 - 0.7238 \ln(x)$  |
| E | $T = 6.25 - 0.5429 \ln(x)$   |
| F | $T = 4.167 - 0.3619 \ln(x)$  |

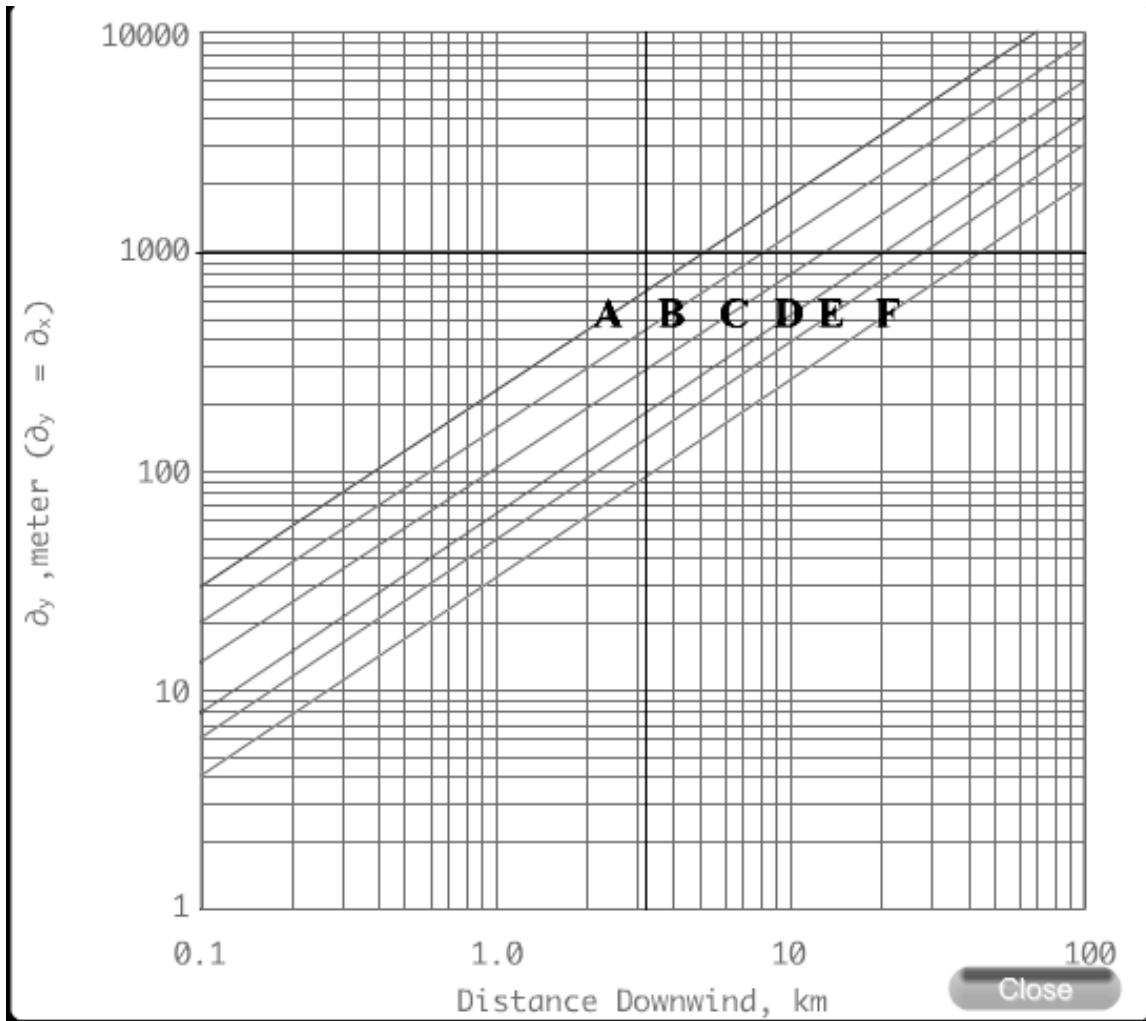


Figure 3. Pasquill-Gifford Horizontal Dispersion Parameters (Turner (1970) and Pasquill (1961))

Dispersion in the vertical (z) direction is somewhat more complicated to predict and again is based on experimental observations. We can estimate the vertical dispersion coefficient,  $\sigma_z$ , by using the same atmospheric stability categories from Table 1 but with a more precise treatment of the wind speed. The equation governing the estimate of vertical dispersion is

$$\sigma_z = a x^b,$$

where  $x$  is the distance in km

$a$  and  $b$  are fitting parameters obtained from Table 3.

A plot of the dependence of vertical dispersion coefficients on distance from the point source is shown in Figure 4. We have been describing dispersion, but what exactly is it? As we have noted, dispersion is a function of the distance from the point source. Dispersion is a mathematical description of mixing between the pollutant plume and the natural atmospheric gases. The values you read from the graph or calculate using the equations are given in meters or kilometers. Thus, the values given represent the width of the pollutant plume at the specified distance from the point source and thus reflect the amount of atmosphere the pollution has mixed with.

Table 3. Pasquill-Gifford Vertical Dispersion Parameter

| Vertical Dispersion Parameter: $\sigma_z = a x^b$ , where $x$ is in km |                   |        |         |                              |
|--|-------------------|--------|---------|------------------------------|
| Stability  | Distance (km)     | a      | b       | $\sigma_z$ at upper boundary |
| A  | > 3.11            |        |         | 5000                         |
|  | 0.5 – 3.11        | 453.85 | 2.1166  |                              |
|  | 0.4 – 0.5         | 346.75 | 1.7283  | 104.7                        |
|  | 0.3 – 0.4         | 258.89 | 1.4094  | 71.2                         |
|  | 0.25 – 0.3        | 217.41 | 1.2644  | 47.4                         |
|  | 0.2 - 0.25        | 179.52 | 1.1262  | 37.7                         |
|  | 0.15 – 0.2        | 170.22 | 1.0932  | 29.3                         |
|  | 0.1 – 0.15        | 158.08 | 1.0542  | 21.4                         |
|  | < 0.1             | 122.8  | 0.9447  | 14.0                         |
| B  | > 0.35            |        |         | 5000                         |
|  | 0.4 - 35          | 109.30 | 1.0971  |                              |
|  | 0.2 – 0.4         | 98.483 | 0.9833  | 40.0                         |
|  | > 0.2             | 90.673 | 0.93198 | 20.2                         |
| C  | all values of $x$ |        | 61.141  | 0.91465                      |
| D  | > 30.             | 44.053 | 0.51179 |                              |
|  | 10. – 30.         | 36.650 | 0.56589 | 251.2                        |
|  | 3 – 10.           | 33.504 | 0.60486 | 134.9                        |
|  | 1 - 3             | 32.093 | 0.64403 | 65.1                         |
|  | 0.3 - 1           | 32.093 | 0.81066 | 32.1                         |
|  | < 0.3             | 34.459 | 0.86974 | 12.1                         |
| E  | > 40.             | 47.618 | 0.29592 |                              |
|  | 20. – 40.         | 35.420 | 0.37615 | 141.9                        |
|  | 10. – 20.         | 26.970 | 0.46713 | 109.3                        |
|  | 4 – 10.           | 24.703 | 0.50527 | 79.1                         |

|   |           |        |         |      |
|---|-----------|--------|---------|------|
|   | 2 - 4     | 22.534 | 0.57154 | 49.8 |
|   | 1 - 2     | 21.628 | 0.63077 | 33.5 |
|   | 0.3 - 1   | 21.628 | 0.75660 | 21.6 |
|   | 0.1 - 0.3 | 23.331 | 0.81956 | 8.7  |
|   | < 0.1     | 24.260 | 0.83660 | 3.5  |
|   |           |        |         |      |
| F | > 60      | 34.219 | 0.21716 |      |
|   | 30. - 60. | 27.074 | 0.27436 | 83.3 |
|   | 15 - 30.  | 22.651 | 0.32681 | 68.8 |
|   | 7 - 15    | 17.836 | 0.4150  | 54.9 |
|   | 3 - 7     | 16.187 | 0.4649  | 40.0 |
|   | 2 - 3     | 14.823 | 0.54503 | 27.0 |
|   | 1 - 2     | 13.953 | 0.63227 | 21.6 |
|   | 0.7 - 1   | 13.953 | 0.68465 | 14.0 |
|   | 0.2 - 0.7 | 14.457 | 0.78407 | 10.9 |
|   | < 0.2     | 15.209 | 0.81558 | 4.1  |

Source: Turner (1970) and Pasquill (1961)

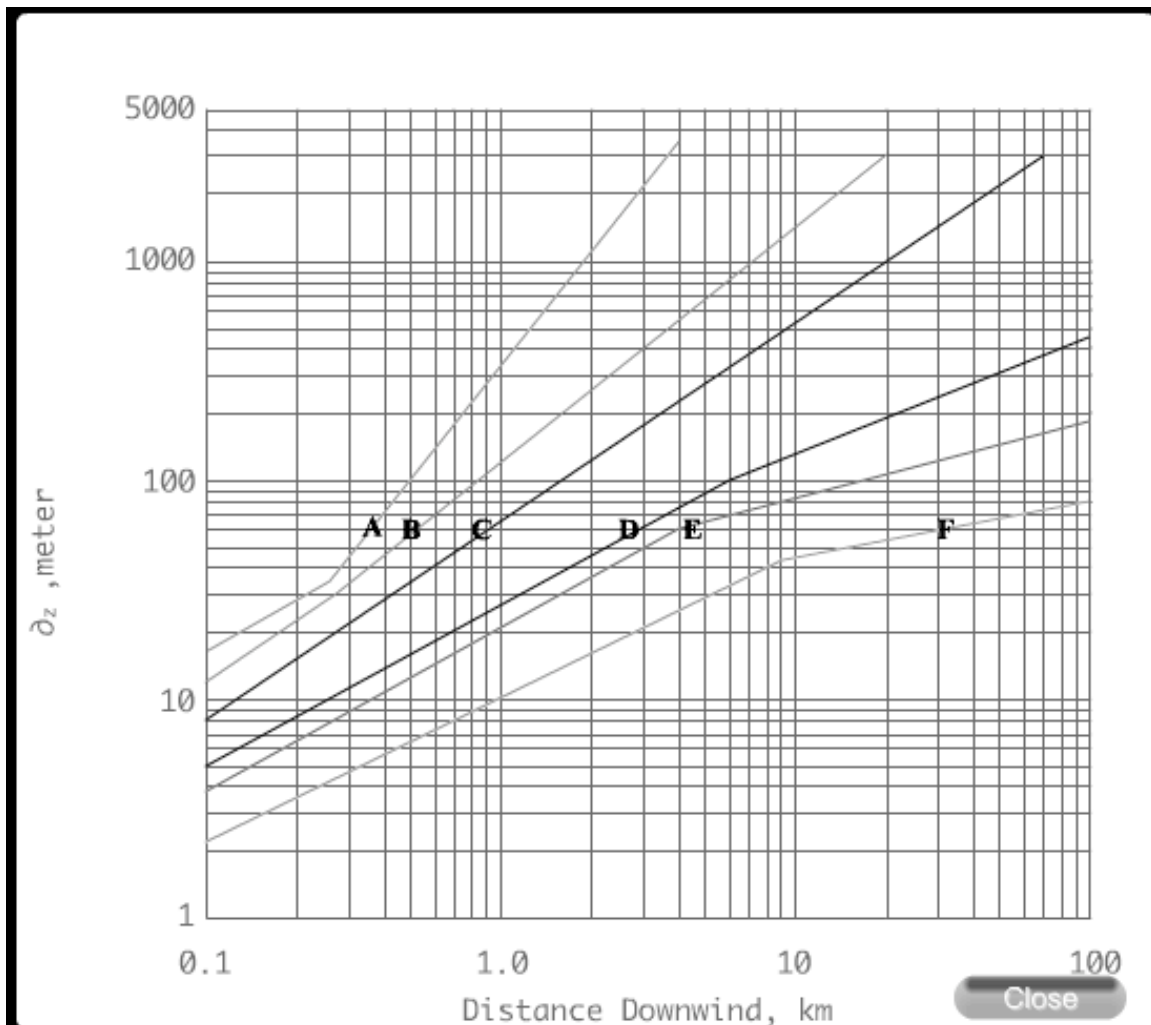


Figure 4. Pasquill-Gifford Vertical Dispersion Parameters (Turner, 1970)

**Pulse Input (Instantaneous or Puff Model) of Pollution:**

For a pulse rather than a plume input, dispersion is handled a little differently. In the step (plume) model we can use either rural or urban dispersion estimates, whereas urban dispersion parameters are usually used for the pulse (puff) model. These dispersion estimates are derived from experimental observations made by McElroy and Pooler (1968) near St. Louis and from Briggs (1972). Calculations for estimating the horizontal and vertical dispersion coefficients are given in Table 4. We will again assume that dispersion in the x and y directions are the same. Atmospheric stability categories are the same as those described in Table 1. Vertical and horizontal dispersion coefficients are shown in Figures 8 and 9.

Table 4a. Urban Dispersion Parameters

| Pasquill Type of Stability | $\sigma_y$<br>(in meters)  | $\sigma_z$<br>(in meters)  |
|----------------------------|----------------------------|----------------------------|
| A – B                      | $0.32 / (0.0004 x)^{-0.5}$ | $0.24 / (0.001 x)^{0.5}$   |
| C                          | $0.22 / (0.0004 x)^{-0.5}$ | $0.20 x$                   |
| D                          | $0.16 / (0.0004 x)^{-0.5}$ | $0.14 / (0.0003 x)^{-0.5}$ |
| E - F                      | $0.11 / (0.0004 x)^{-0.5}$ | $0.08 / (0.0015 x)^{-0.5}$ |

For distances, x, between 100 and 10000 meters

Source: Turner (1994), Briggs (1972), and McElroy and Pooler (1968)

Table 4b. Open-Country Dispersion Parameters (NOT used in Fate<sup>®</sup>, but you may manually enter the calculated values)

| Pasquill Type of Stability | $\sigma_y$<br>(in meters)       | $\sigma_z$<br>(in meters)       |
|----------------------------|---------------------------------|---------------------------------|
| A                          | $0.22 x / (1 + 0.0001 x)^{0.5}$ | $0.20 x$                        |
| B                          | $0.16 x / (0.0001 x)^{0.5}$     | $0.12 x$                        |
| C                          | $0.11 x / (1 + 0.0001 x)^{0.5}$ | $0.08 x / (1 + 0.0002 x)^{0.5}$ |
| D                          | $0.08 x / (1 + 0.0001 x)^{0.5}$ | $0.06 x / (1 + 0.0015 x)^{0.5}$ |
| E                          | $0.06 x / (1 + 0.0001 x)^{0.5}$ | $0.03 x / (1 + 0.0003 x)$       |
| F                          | $0.04 x / (1 + 0.0001 x)^{0.5}$ | $0.016 x / (1 + 0.0003 x)$      |

For distances x between 100 and 10000 meters

Source: Turner (1994), Briggs (1972), and McElroy and Pooler (1968)



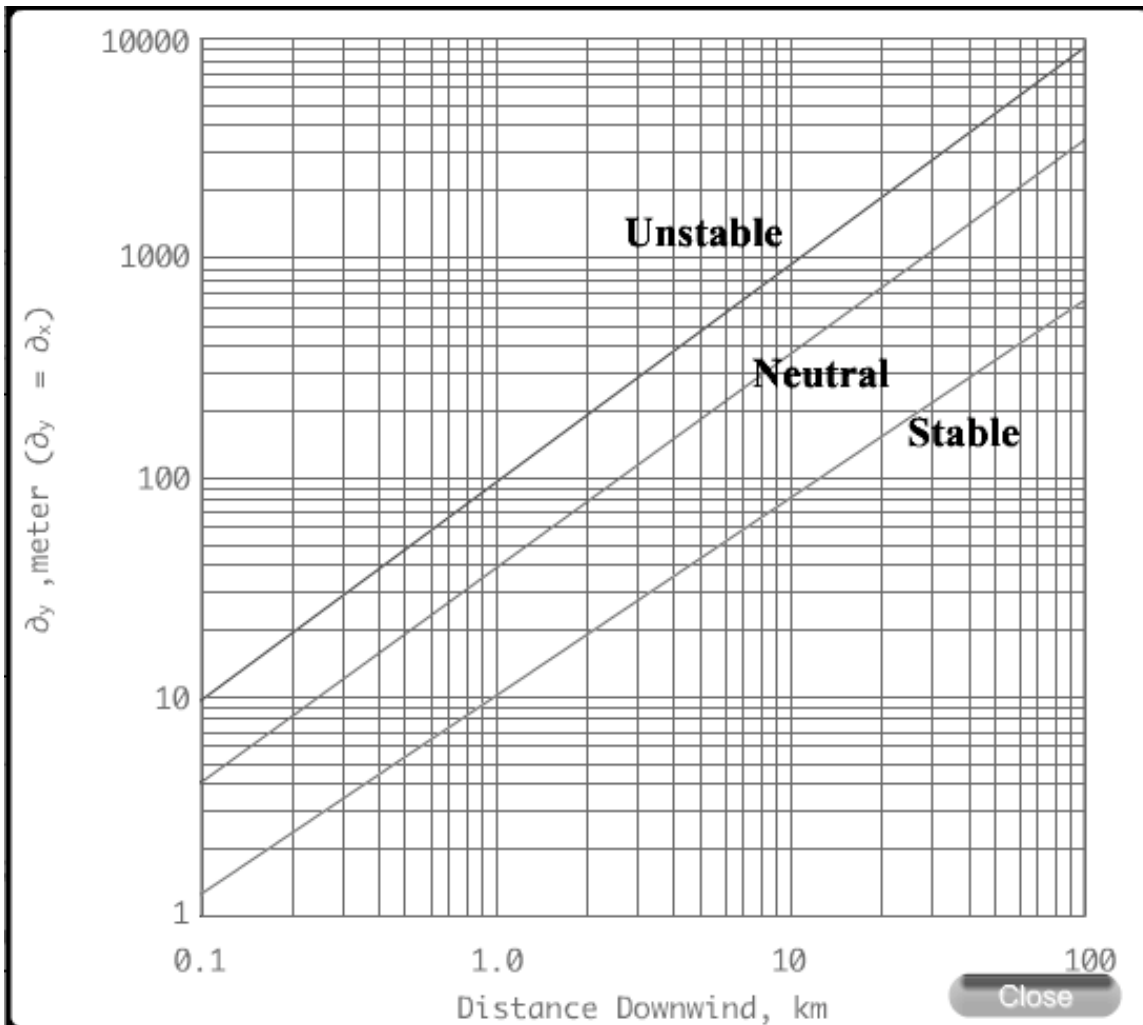


Figure 8. Pasquill-Gifford Vertical Dispersion Parameters (Turner, 1970)

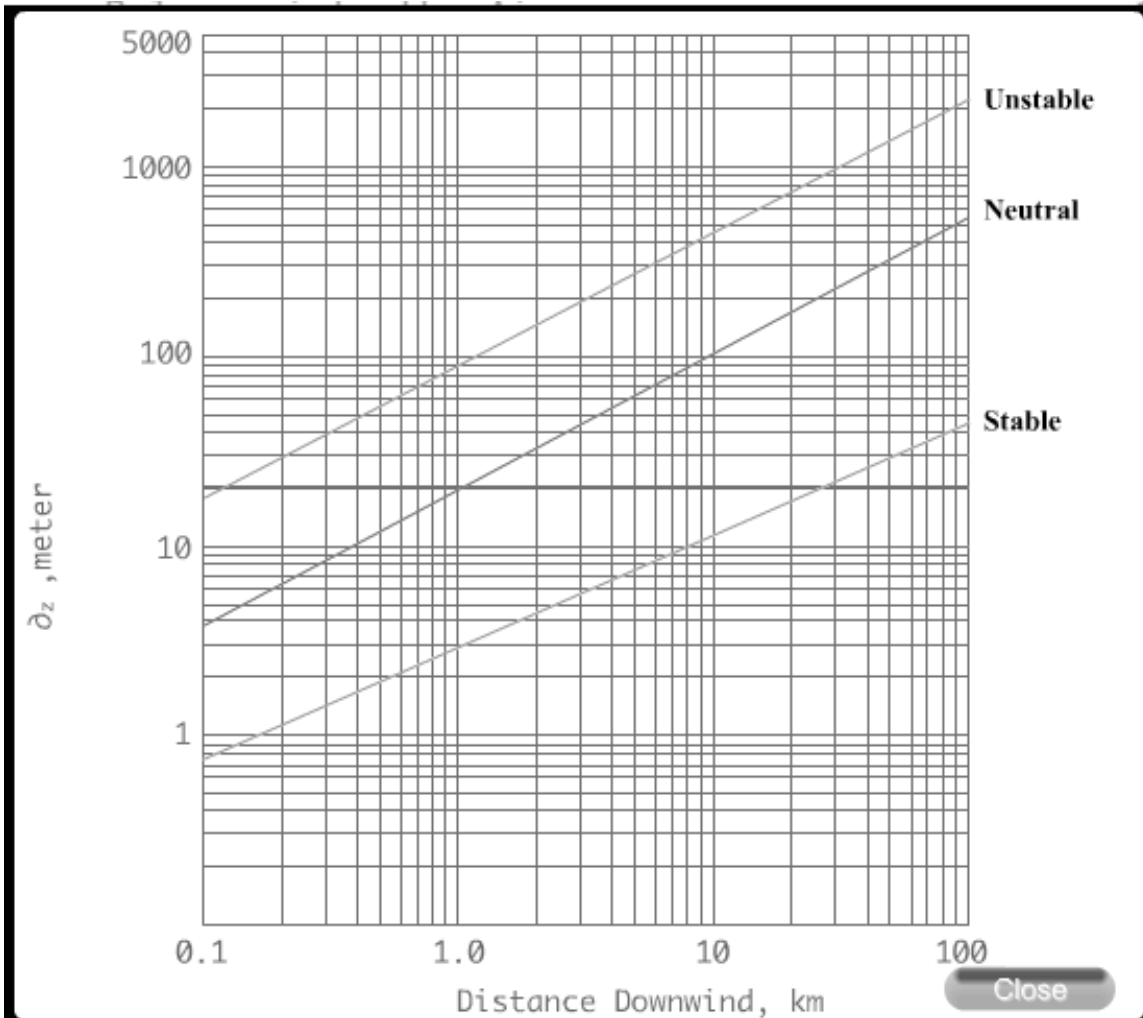


Figure 9. Pasquill-Gifford Vertical Dispersion Parameters (Turner, 1970)

Using the stability categories, wind speed, and the equations shown in Table 4, we can now estimate the atmospheric pollutant concentration downwind from an instantaneous (also referred to as pulse or puff) source by

$$C(x,y,z,t) = \frac{Q_m}{(2\pi)^{3/2} \sigma_x \sigma_y \sigma_z} \exp\left[-\frac{1}{2} \frac{y^2}{\sigma_y^2}\right] \exp\left[-\frac{1}{2} \frac{(z - H_r)^2}{\sigma_z^2}\right] + \exp\left[-\frac{1}{2} \frac{(z + H_r)^2}{\sigma_z^2}\right] \quad \text{Eqn}$$

where  $C(x,y,z, t)$  is the concentration of pollutant in the plume as a function of  $x$ ,  $y$ , and  $z$  (mass/length<sup>3</sup>),  
 $x$ ,  $y$ , and  $z$  are distances from the source (length) see Figures 3 and 4,  
 $t$  is time  
 $Q_m$  is the pollutant source (mass/time)  
 $\sigma_x$  and  $\sigma_y$  are the horizontal dispersion coefficients (length),  
 $\sigma_z$  is the vertical dispersion coefficient (length),  
 $H_r$  is the height of the release (length).

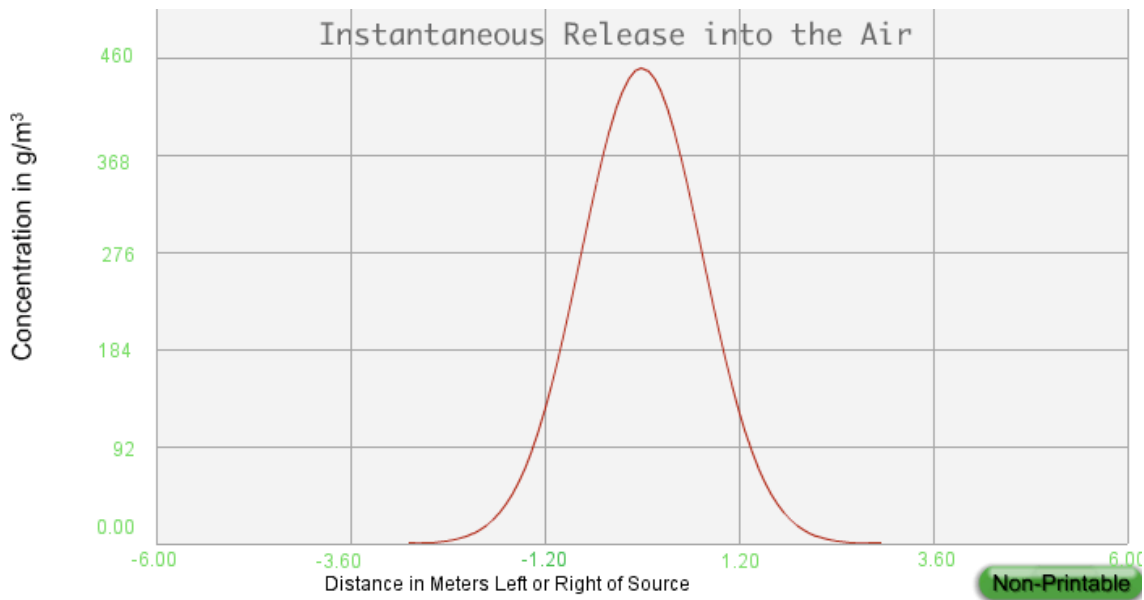
Note the inclusion of time since the distance traveled (x) is a function of wind velocity (u) and time (t), where

$$x = ut.$$

For the concentration along the centerline (y = 0, z = 0, and H<sub>r</sub> = 0) we can use a simplification of equation 3, to yield

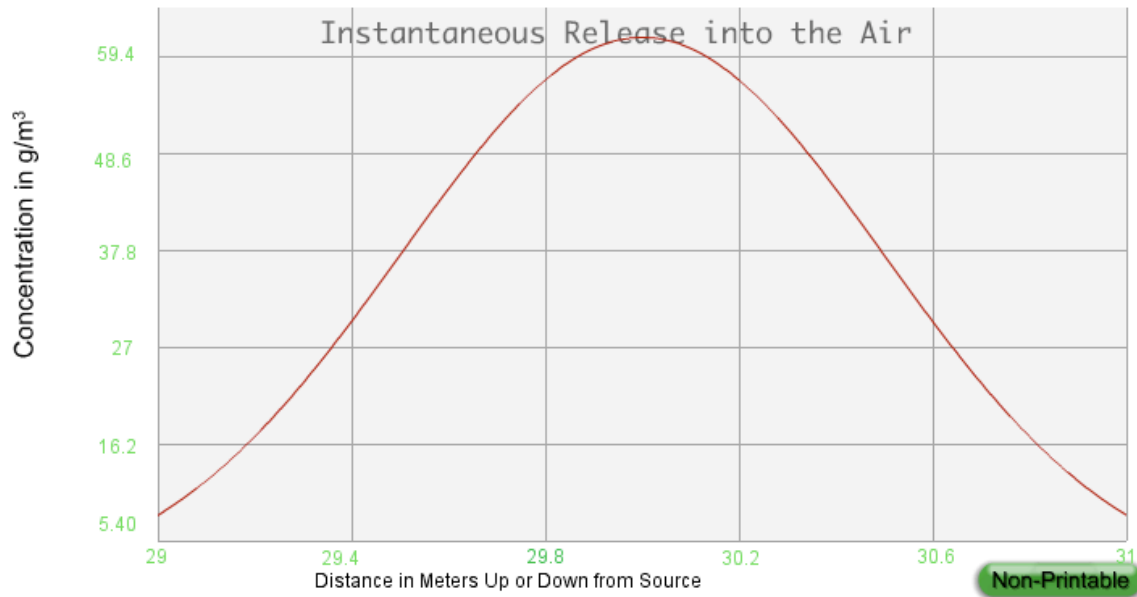
$$C(x,0,0,t) = C(ut, 0, 0, t) = \frac{Q_m}{(\sqrt{2\pi})^{3/2} \sigma_x \sigma_y \sigma_z} \quad \text{Eqn 4}$$

Simulation output from Fate<sup>®</sup> are shown in Figures 10 and 11 for viewing pollutant concentration along the y-axis and the z-axis, respectively.



|             |                             |                           |
|-------------|-----------------------------|---------------------------|
| X = 10 km   | Q <sub>m</sub> = 2000 g/Sec | σ <sub>x</sub> = 0.7517 m |
| Z = 30 m    | ū = 4 m/Sec                 |                           |
| T = 2.5 Sec | H <sub>r</sub> = 30 m       | σ <sub>z</sub> = 0.4990 m |

Figure 10. Output from Fate<sup>®</sup> for a pulse release (puff) of pollutant into the atmosphere with variation in horizontal distance from source.



|             |                     |                       |
|-------------|---------------------|-----------------------|
| X = 10 km   | $Q_m = 2000$ g/Sec  | $\sigma_x = 0.7517$ m |
| Y = 1.5 m   | $\bar{u} = 4$ m/Sec |                       |
| T = 2.5 Sec | $H_r = 30$ m        | $\sigma_z = 0.4990$ m |

Figure 11. Output from Fate<sup>®</sup> for a pulse release (puff) of pollutant into the atmosphere with variation in vertical distance from source.

### References:

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